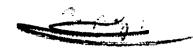
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TECHNICAL NOTES



NATIONAL ADVISORY COMMITTEE FOR AERONAUTICS

No. 439

METEOROLOGICAL CONDITIONS DURING THE FORMATION

OF ICE ON AIRCRAFT

By L. T. Samuels Weather Bureau

Ta he returned to

the files of the Longley Memorial Aeronautical Laboratory.

Washington December, 1932



NATIONAL ADVISORY COMMITTEE FOR AERONAUTICS

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OF ICE ON AIRCRAFT

By L. T. Samuels

INTRODUCTION

The hazard resulting from the formation of ice on airplanes makes it highly desirable to ascertain all possible meteorological information pertinent thereto in order to avoid or prevent its formation. The object of this paper is to present the results of a number of records recently secured from autographic meteorological instruments mounted on airplanes at times when ice formed.

Ice is found to collect on an airplane in appreciable amounts only when the airplane is in some form of visible moisture, such as cloud, fog, mist, rain, etc., and the air temperature is within certain critical limits.

There are two principal types of ice formation that collect under such conditions, and in view of the different effects of each of these on a plane they will be discussed separately under their respective headings, viz., clear ice and rime.

Clear ice. This is the same type as that commonly known as "glaze," which forms on the ground, trees, and other objects from rain when the temperature of these objects is below 0°C. It usually is smooth and glassy in appearance, but when mixed with snow or sleet it may be rough; also, when freezing takes place slowly, ridges are likely to form.

This deposit usually is heaviest on the entering edge of the plans where it assumes a blunt-nosed shape tapering off toward the rear. Occasionally the wings are ice-coated on both top and bottom with icicles along the trailing edge. In most cases the ice adheres firmly to the surface of the airplane.

Rime. - Rime consists of hard, whitish, opaque ice pellets or grains, frequently intermixed with a frost formation of light feathery crystalline structure. From observitions of rime deposits on mountains, Köhler (reference 1) described the formation as snow-white, plug-like, truncated of the small end toward the surface upon which it is deposited. The plugs showed a fibrous structure and occasionally shiny surfaces. The particles from which the plugs were composed were firmly held together but the plugs themselves could easily be separated from one another. Their interior was usually of granular appearance. The spaces between the plugs were filled with a powder composed of these grains. From laboratory tests by Scott (reference 2) the granular structure of rime appeared to be coarsor at the lower temperatures of formation.

Unlike clear ice, rime builds outward from the leading edges of the airplane into a sharp-nosed shape. As a rule it does not adhere to the plane as firmly as clear ice and is less resistant to the vibration and wind force encountered in flight.

Frost. - A third type of deposit of lesser importance, however, than clear ice and rime, which sometimes forms on airplanes is frost. This is of a light feathery crystalline structure such as often is observed on ground objects in the early morning. It does not adhere to the airplane very firmly and is never dangerous as it has very little resistance to the vibration and wind force encount tered in flight.

Effects of ice deposit. As a rule the first notice—able effect of an ice deposit is an increase in vibration of the airplane followed by increasing difficulty in its control. As the deposit becomes heavier the vibrations may cause severe structural strains with a possibility of fracturing individual parts. The deposit frequently stops up the nozzle of the air-speed indicator thus rendering that instrument useless. Other instruments may also be affected. An ice formation on the propeller is likely to produce a difference in weight of the blades which may become sufficient to cause the engine to break loose. Forced landings frequently are necessary due to icing.

The chief dangers to lighter-than-air craft result from the ice being thrown off the propeller and possibly puncturing the gas containers. Also, the distribution of the ice on the airship may cause structural strains with

the possible collapse of a nonrigid or semirigid type. Ice collections on the radio antennae may result in their breaking. In general, however, ice deposits on lighter-than-air craft are less serious in their effects than on airplanes. (Reference 3.)

A deposit of clear ice is a greater hazard to an airplane in flight than one of rime. The reason for this is that clear ice formations, by virtue of their blunt-nosed shape, break up the normal air flow over the surface of the airplane and thereby reduce the lift, increase the friction, and cause excessive vibration. The weight of the ice also adds to the danger, although this factor in itself usually is of less importance. On the other hand, the contours formed by rime produce less detrimental aerodynamical effects, and moreover, rime is much more easily blown and shaken off the airplane.

Results of observations.— The records obtained by the Weather Bureau, previously referred to, were classified according to the two general types of formation, viz., clear ice and rime, together with the respective temperatures, relative humidities, clouds, and elevations above ground at which the formations occurred. This classification includes 108 cases where rime formed, 43 cases in which clear ice formed, and 4 cases when both rime and clear ice formed during the same flight, It is evident from the above figures that there was a preponderance of rime by the ratio of 2.5 to 1, while in only a very few cases both types of ice formation occurred during the same flight.

Table II contains a summary of the observations shown in Table I. In examining these tables, it should be kept in mind that the airplane usually continued to climb after ice began to form and therefore the temperature was generally lower where the formation ceased than where it began. Also, the heavier coatings are in most cases a consequence of the airplane being subjected to ice-forming conditions for a longer time than when the lighter coatings formed.

The following points of interest are brought out in Table II:

1. The temperature averaged 1.8°C, lower during the rime formations than during those of clear ice.

2. The average temperature interval, i.e., range of temperature from beginning to end of formation. for rime and clear ice was the same, viz., 3.1°C. ・ 17 (基準型・ 1.11.g → 1.11)

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- 3. The relative humidity averaged practically the same during the rime formations as during those of clear ice. In this connection it should be stated that owing to the fact that the airplane usually was climbing while the ice formed, the resulting lag in the humidity element would tend to indicate values somewhat too low, particularly at the lower part of the stratum wherein the ice formed. However, it is reasonable to assume that unsaturated conditions frequently obtain within clouds since tabulations by Pick (reference 4) show that fogs often occur with a relative humidity of less than 100 per cent. The second secon
- 4. The average relative humidity interval, i.e., the range of humidity from beginning to end of formation, was practically the same for rime and clear ice, the humidity being about 2 per cent higher at the end of the formation than at the beginning. This difference is probably due to a large extent to the lag mentioned above in paragraph 3.
- 5. The average time interval during which the formations occurred was only slightly greater for clear ice than for rime, viz. 6.3 and 6.1 minutes, respectively.
- 6. The average elevations at which rime formed were somewhat higher than those for clear ice, with the exception of the heaviest coatings, in which case the clear ice formations occurred at a greater average height than those of rime.
- 7. The average thickness of the stratum in which the formation occurred was somewhat greater for clear ice than Applications of the second section of the section for rime.
- 8. The fact that clear ice has, in general, a more serious effect than rime is well brought out by the following figures which give the percentage of flights listed in Table I which were terminated because of ice formation.

	Terminated because of ice formation	Thickness of ice
Clear ice	31 per cent	1/4 inch or more
Clear ice	12 "	Less than 1/4 inch
Rime	4 "	1/4 inch or more
Rime	0 "	Less than 1/4 inch

In Figure 1 are shown the actual number of cases when each of the two types of ice formation occurred at various temperatures, the latter representing the mean of the temperature range through which the formation occurred.

Probably the most outstanding feature shown by this chart is the fact that rime formed more frequently than clear ice at <u>all</u> temperatures. From this it is obvious that temperature alone cannot be used as a safe criterion for indicating which type of formation will occur on any particular occasion.

Additional features shown are:

- 1) The temperatures at which the most frequent deposits occurred were higher for clear ice $(-4^{\circ}\text{C}. \text{ to } -5^{\circ}\text{C}.)$ than for rime $(-6^{\circ}\text{C}. \text{ to } -7^{\circ}\text{C}.)$. These values agree closely with those found by Peppler (reference 5) from kite observations. The latter indicated that clear ice formed at an average temperature of -4°C . and rime at -6°C .
- 2) The extreme temperature range during clear ice deposits $(0.5^{\circ}\text{C}. \text{ to } -17.5^{\circ}\text{C}.)$ was slightly less than that for rime $(0.5^{\circ}\text{C}. \text{ to } -20.5^{\circ}\text{C}.)$.

In this connection it is interesting to note that Köhler (reference 1) observed a thin layer of clear ice which formed from a "wet fog" at -23.6° C.

In Figure 2 are shown the percentage frequencies of clear ice and rime formations at various temperatures, the latter, as in Figure 1, representing the mean of the temperature range through which the formation occurred. It should be understood that the percentages indicated in Figure 2, as well as those in Figures 3, 4, and 5 are with reference to the total number of each of the respective types and not to the total number of both types.

A significant feature shown in this chart is the tendency for clear ice to form at relatively higher temperatures a greater percentage of the time than rime. By computation it is found that 58 per cent of the total number of clear ice formations occurred at temperatures at, or above, -5°C., whereas only 37 per cent of the total number of rime formations occurred at those temperatures.

Figure 3 shows the percentage frequencies of both types of formation at the temperatures at which the ice began to form.

It is evident from this chart that the clear ice formations began most frequently at a slightly higher temperature (0°C.) than those of rime (-2°C.). Both types began forming at the same maximum temperature (1°C.). Rime began forming at a lower temperature (-20°C.) than clear ice (-17°C.). By computation it is found that in 74 per cent of the total number of clear-ice deposits the temperature at the beginning of the formation was -5°C., or higher, as compared to only 55 per cent in the case of rime.

Since, as previously pointed out, both clear ice and rime may form over practically the same range of temperature it is necessary to conclude that one or more other factors are decisive in determining which type is deposited. We may also conclude that the other decisive factors just referred to change in their potency or relative frequency with temperature, so that the factors favoring the formation of clear ice, for example, are more frequent or more powerful, or both, at higher temperatures than at lower temperatures.

In Figure 4 are shown the percentage frequencies of clear ice and rime formations at various heights above the ground, the latter representing the mean of the height interval in which the formations occurred. The following features are evident in this chart:

- 1) A very pronounced maximum frequency of occurrence of both clear ice and rime at relatively low heights, viz., between 500 and 1,000 m (1,640 and 3,281 ft.).
- 2) Pronounced secondary maximum frequencies of occurrence between 2,500 and 3,000 m (8,200 and 9,842 ft.) for clear ice and between 4,000 and 4,500 m (13,123 and 14,764 ft.) for rime. These primary and secondary maximum frequencies of occurrence are possibly related to layers of maximum condensation. Such a layer was found by Lewis (reference 6) between 500 and 1,000 m (1,640 and 3,281 ft.) above ground where the maximum frequency of strato-cumulus cloud bases occurred.
- 3) Low frequencies of occurrence of both types of ice formation between 1,500 and 2,500 m (4,920 and 8,200 ft.) above ground.
- 4) Both types of formation occurred throughout the same strata and with small and practically equal percentages of frequency at the lowest and greatest heights

reached. The maximum heights where icing occurred coincide with the maximum heights of the flights.

In Figure 5 are shown the percentage frequencies of both types of ice formation for various cloud and weather conditions. The following features are brought out:

- 1) Both clear ice and rime formed most frequently in strato-cumulus clouds.
- 2) When in rain but not in cloud, the formation was always clear ice, whereas when in rain and cloud, the formation was sometimes rime and sometimes clear ice.
- 3) When above cloud and not in any form of precipitation, the formation was always clear ice. In such cases the deposit formed from moisture collected on the airplane while passing through the cloud.
- 4) Comparatively high percentage frequencies of clearice formations occurred in alto-stratus clouds and of rime in stratus clouds.
- 5) No ice deposits were reported in cumulus clouds. This is doubtless due to the fact that most of the observations were made before daylight (about 5 a.m., 75th meridian time), when cumulus clouds are seldom present.

An examination of the prevailing temperature lapse rates occurring in these observations showed no relation-ship between the lapse rates and the types of ice formation.

The relative distribution of the number of occurrences of both types of formation from the data at hand is shown in Table III. It will be noted therein that the ratios of rime and clear ice deposits vary considerably for the four stations. As was previously stated this ratio for the observations for all stations combined was 2.5 to 1, with a preponderance of rime. However, these ratios for the individual stations are as follows: Chicago, 1.7; Cleveland, 7.5; Dallas, 5.5; and Omaha, 0.6. It is also found that the ratios between the light and heavy deposits vary considerably among the individual stations, e.g., the heavy coatings of clear ice predominate at Chicago and Cleveland, whereas the light coatings of clear ice predominate at Omaha and apparently at Dallas. The heavy coatings of rime predominate at Cleveland and apparently at Dallas, whereas the light coatings of rime predominate at Chicago and Omaha.

Cleveland had the greatest number of rime formations, with Chicago second, Dallas third, and Omaha fourth. This same order, however, did not occur in the case of clear ice. While this may be partly due to the smaller number of observations of this type, it is probably due also in part to other factors such as available nuclei and prevailing winds with respect to nearby water areas.

It is shown in Table IV that the average temperature was lower during the rime formations than during those of clear ice at every station.

A comparison of the ratios of the total number of cases of both types of ice deposits to the average amount of lower clouds reveals no proportionality. (See Tables III and IV.) The average heights where the formations occurred were approximately the same as the average heights of strato-cumulus (i.e., lower) clouds in which the maximum number of formations occurred. (See Table II.) Therefore other conditions than the incidence of clouds at subfreezing temperatures must be sought as controlling factors. A possibility in this connection is the relative number of available nuclei as a factor in the determination of the size of the cloud droplets.

A comparison of the average temperatures during the ice formations (Table IV) with the average for the season at corresponding heights shows lower temperatures during the times of formation with one exception, viz., Cleveland, for clear ice. This station had relatively few cases of clear ice deposits and it seems probable that this relatively high average temperature at the time of formation is due to the proximity of Lake Erie and the prevailing winds which were mostly off the lake at those times. Greater temperature differences will be noted in the case of rime than for clear ice at all stations.

Factors bearing on the type of ice formation. It has been shown that other factors than temperature have an important bearing on the nature of the ice deposit, i.e., clear ice or rime. One of these factors presumably is the size of the water droplets. It seems probable that, in general, large droplets tend to form clear ice, whereas small droplets usually produce rime. This view is strengthened by the fact that the deposit formed while flying in rain, i.e., when not encountered with cloud particles, is always of the clear ice type. Köhler (reference 1) came to the conclusion that when a sufficient number of large un-

dercooled droplets impinge on a suitable object, the freezing of a portion of the water deposited liberates latent heat of fusion which, if not conducted away with sufficient rapidity, causes the temperature of the deposit to rise, possibly as high as 0°C. This permits the spreading and flowing of the water droplets referred to above and a layer of liquid admixed with some ice results. By virtue of this higher temperature the saturation vapor pressure over the deposit will now be higher than the vapor pressure about the subcooled droplets in the cloud and evaporation will occur and hence a cooling of the deposit, an effect which under the conditions given above, when combined with the loss of heat by conduction to the passing air stream and to objects upon which the water is deposited produces freezing of the remaining liquid and gives rise to clear ice.

On the other hand, small droplets are more likely to freeze immediately upon striking the airplane. This is in part due to the greater convexity and different distribution of mass and cohesive forces in smaller droplets, all of which hinder them from spreading and flowing and aid in maintaining their spherical form. Since there is a greater exposed surface area about a given mass of water in the form of small droplets than about an equal mass which has spread and flowed from larger droplets, the removal of the latent heat of fusion liberated is probably more rapid in the former case. Hence, in general, small droplets have a greater speed of crystallization than have large droplets, a condition which, in the opinion of Kohler (reference 7), is conducive to the formation of rime.

It seems probable, however, that small droplets might also produce clear ice where the circumstances are such that the liberated heat of fusion is not conducted away with sufficient rapidity.

Köhler thought it probable that the type of ice formation depends to a considerable extent upon the speed of crystallization at which the liquid water freezes, there being a critical value for this speed which, when exceeded, produces rime or frost and when unattained produces clear ice. He thought it possible also that a higher critical value of the speed in question might exist which, when exceeded, produced frost instead of rime.

The speed of crystallization, in turn, depends on the degree of concentration of the dissolved salts serving as nuclei and on the temperature of the subcooled droplets.

(Reference 7.) Thus for a given concentration and a relatively low temperature the speed of crystallization is relatively high, whereas, for the same concentration at a relatively high temperature the speed of crystallization is relatively kow. Also, for a given temperature and a low concentration the speed of crystallization is relatively high, whereas, at the same temperature and a high concentration the speed of crystallization is relatively low.

From measurements of the concentration of salts in clear ice and rime deposits and the corresponding sizes of fog and cloud droplets on mountains in Europe, together with certain assumptions, Köhler (reference 7) concludes that the sizes of droplets in clouds, from which no precipitation is falling and which exist simultaneously at the same elevation, depend on the respective sizes of the salt particles about which condensation has occurred. His calculations show that high concentrations are associated with small droplets and, vice versa.

It will be noted that it was stated above that small droplets are associated with a high concentration of salt nuclei, that the latter produces a relatively low speed of crystallization and further, that the latter generally tends to produce clear ice. From other considerations it was concluded that small droplets generally tend to form rime. Thus we find from two sets of considerations that small droplets tend to form both clear ice and rime. It must therefore be concluded that further investigation of this phase of the subject is necessary in order to determine qualitatively and quantitatively the manner in which the various factors operate to produce the particular type of ice deposit. A parallel line of reasoning applies to large droplets.

Another possible factor bearing on the type of ice deposit is the mass of water striking a unit area in unit time. (Reference 7.) It is obvious that the mass of water in question depends on the amount of water per unit volume of the cloud and on the speed of the airplane. When the mass of water striking a unit area in unit time is large a sufficient amount of latent heat may be liberated so as to produce clear ice in the manner previously described.

Scott found from wind-tunnel experiments that the air speed apparently has little effect upon the character

of the ice formation. (Reference 2.)

The formation of frost on aircraft, previously referred to, is a result of sublimation, i.e., a change directly from the gaseous to the solid state, and therefore requires a state of supersaturation with respect to ice.

Sudden ice deposits.— It has been suggested by various authors that supersaturation, with respect to ice, in clouds composed of subcooled water droplets may be responsible for comparatively sudden and heavy deposits occasionally reported by pilots. Humphreys (reference 8) has shown, however, that at a temperature of -10°C., if all of the excess vapor in the air, i.e., assuming a condition of supersaturation with respect to ice, were deposited, it would be equivalent to a layer of clear ice one inch thick on the front of an airplane after the latter had flown for a distance of 72 miles. It is probable though that only a small part of the excess vapor encountered would be deposited on the airplane

An occurrence of a sudden deposit together with a possible explanation was recently reported by A. Hansen. (Reference 9.) The following is quoted therefrom.*

"In a summer cumulus cloud with strong heat convection, the speed indicator stopped functioning almost immediately because of icing of the nozzle upon flying into the cloud. The bumpiness was such that the airplane did not respond to the movements of the rudder. After 5 or 10 seconds the corrugated ribs on the top side of the wing were concealed under a layer of ice, which had not thickened on the front edges, but the entire visible wing surface was apparently equally heavily coated. The thermometer showed about 000., the air was very wet, the height was about 3,600 m (11,811 ft.). In consequence of the excessive demands, the airplane quickly lost altitude in spite of the thermal convection and wide-open engine and soon fell out of the cloud base. Here the ice melted quickly and at about 1,000 m (3,281 ft.) had completely disappeared.

"The suddenness of the icing and the unusual form of the ice cover even in the region of dynamic pressure reduction cannot be explained in the usual manner. It seems

^{*}Translated by J. C. Ballard, Aerological Division, Weather Bureau.

plausible that here the "triple point"* plays a part. If the vapor pressures over water and ice are equal, the heat of vaporization and heat of fusion may be exchanged for one another in the presence of liquid water. Since the process is intermolecular and no external heat exchange is assumed, it can take place practically instantaneously. Since the heat of vaporization is about eight times as great as the heat of fusion, a partial evaporation must form an eightfold quantity of ice. The vaporization can be caused by dynamic pressure reduction on the airplane; for example, on the top side of the wing. Through consideration of the triple point, the manner and speed of this special type of icing follow quite naturally."

In connection with the foregoing, it is interesting to note in Table I that in most cases where rapid icing occurred the temperature was not much below the freezing point, and it seems possible that the physical explanation of at least a part of the ice formation in those cases is similar to that given by Hansen.

In general, no ice formation will occur at temperatures above freezing. However, occasionally cases are reported where it does form in wet clouds or in rain at temperatures slightly above freezing, and in such cases it is probable that the ice is formed by evaporative cooling, the extent of which varies inversely as the relative humidity.

Undercooled water droplets. - In connection with the occurrence of undercooled cloud droplets, it is of interest to note that these are found at surprisingly low temperatures. A. Wegener (reference 10) observed a "fog-bow" in Greenland at a temperature of -34°C., indicating that the fog particles were in the liquid state.

At Little America, headquarters of the Byrd Antarctic Expedition, both cloud and fog particles were frequently observed in the liquid state at very low temperatures. W. C. Haines (reference 11), meteorologist of this expedition, states as follows regarding this:

^{*}The triple point is the temperature (0.0072°C.) and vapor pressure (4.58 mm of mercury) for which the three states - vapor, liquid, and solid - can exist together in equilibrium. At the triple point the saturated vapor pressures for ice and water are identical.

"Fog, while infrequent, was interesting from the point of view of showing that water particles can, and do, exist in the atmosphere at temperatures far below the freezing point. Considerable attention was given by Mr. Henry T. Harrison and myself in observing this phenomenon. We used great care in examining fog or mist when it occurred before recording it as such. In every case when the fog was dense and lasted for an appreciable length of time, a deposit of rime would form on the windward side of objects due to the impingement of the undercooled fog particles. Fogs were observed at temperatures of -26°C., -30°C., and -44°C.

"During kite flights at Little America, when clouds of the stratus or strato-cumulus type were entered, the kites and wire would always be covered with rime on reeling in, thus proving beyond doubt that the clouds were composed of water particles. The lowest temperature observed at the cloud base was approximately -18°C. However, these clouds had the same appearance as those of similar type observed at -45°C., or -50°C. Who can say definitely but that they also were composed of water particles?"

The complete explanation of the manner in which water exists in the liquid state at such low temperatures is not known. Köhler (reference 7), from his investigation of the solid substances found in rime, ice, and snow, is of the opinion that this is primarily due to the concentration of salts dissolved in the droplets. While Köhler is inclined to believe that sea salt is the chief source of these nuclei, Lenard and Ramsauer (reference 12) have shown that the effects of ultra-violet solar radiation upon certain atmospheric constituents may produce hygroscopic nuclei of composition different from that of sea salts and equally, or more effective in respect to their hygroscopic properties. Such substances dissolved in the droplets would have the same effect as regards subcooling as seasalt nuclei. Solutions of any of these substances may be cooled to various temperatures below 0°C., before freezing occurs, depending on the concentration and kind of substance, the degree of ionization, the radius of the droplets and possibly, also, on other factors.

Other conditions favorable or unfavorable to ice formation: A deposit of frost may occur when an airplane descends rapidly from a region where the temperature is below freezing into a warmer, but still subfreezing stratum, which is nearly or entirely saturated. In such a case the formation occurs instantly but stops as soon as the airplane attains the same temperature as that of the surrounding air.

Another condition conducive to frost is in air nearly saturated and the temperature at, or below freezing. The reduced air pressure, and consequently lowered temperature, just above the wings in such a region might be such that condensation would cause a small amount of frost to form.

Sleet, by itself, does not collect on an airplane. However, when mixed with rain it is likely to form a rough and dangerous coating.

Clouds composed of ice spicules do not form any appreciable deposit.

Dry snow does not adhere to an airplane. A mixture of snow and rain or cloud droplets, however, is likely to form a dangerous deposit of frozen slush.

Ice deposits from freezing rain may often be partially removed or prevented by flying in the inversion, i.e., warmer layer, which usually exists above such rains.

A light deposit may form on an airplane flying in a region where cloud droplets are of such small size as to render them invisible, providing the temperature is below freezing.

Methods of determining whether ice will form. The fact that ice deposits of appreciable amounts do not occur unless the airplane is in some form of visible moisture, is of prime importance because in this way the pilot is visually warned, providing he knows the air temperature. The latter can be ascertained by means of a distant indicating thermometer. At night visible moisture can generally be detected by means of a light on the airplane.

In view of the important difference in the effects of clear ice and rime formations on an airplane in flight, it is obvious that any means of determining which of the two types is likely to form on any particular occasion would be of great benefit. While temperature cannot be used as a sole criterion in regard to the particular type of ice formation it is, however, the principal criterion as regards the probability of any formation at all. With-

out upper air observations the temperature aloft must, of course, be estimated from surface conditions. To do this properly one must assume a certain lapse rate, i.e., vertical change in temperature. The lapse rate prevailing at any particular time depends on a variety of factors, the principal ones being (a) time of day, (b) season, (c) latitude, (d) nature of the surface, i.e., land or water, (e) cloudiness, (f) wind velocity and direction, (g) atmospheric pressure distribution, and (h) precipitation. As many of these factors as possible should be taken into consideration.

The following will assist in estimating the temperature lapse rate at any particular time.

On the average the temperature decreases about 0.6°C. per 100 m (328 ft.) elevation. In the lower levels, i.e., the first 1.000 m (3.281 ft.) or so, the lapse rate may vary from slightly more than 1°C. per 100 meters to a large negative value, i.e., the temperature may increase with elevation. The latter condition is called a temperature inversion and is a common phenomenon at night and early morning during clear, calm weather. It is most pronounced in winter and at higher latitudes. A solid cloud layer at night tends to minimize the intensity of the nocturnal inversion, as then terrestrial loss of temperature is materially reduced by the return radiation from the cloud. The intensity of nocturnal inversions is likewise reduced by wind which mixes the air and thereby prevents extreme stratification.

During mid-afternoon, particularly in the warmer season, the lapse rate generally increases until it reaches, or slightly exceeds, the adiabatic rate for dry air, i.e., 1°C. per 100 meters. An overcast sky during the daytime tends to keep the lapse rate low, as then the clouds intercept a large part of the solar radiation by absorption and reflection.

Precipitation tends to decrease the lapse rate. The lapse rate within a cloud is usually less than in clear air, except that immediately above sheet clouds there is often a temperature inversion.

For more detailed information regarding the effects of these and other factors bearing on the temperature lapse rate reference should be made to a good textbook on meteorology.

The determination of the size of cloud droplets from ground observations is difficult. An incipient rain condition is a fairly certain indication of large droplets. This may be indicated by the appearance of the clouds and a knowledge of general weather conditions at surrounding stations.

Some indication of the size of cloud droplets is afforded by the presence of a corona which, if very close to the sun, or moon, signifies relatively large droplets, whereas, when the corona is large, i.e., farther away from the sun, or moon, the droplets are correspondingly smaller. This criterion, however, would probably be of little practical value, since coronae are visible only when the clouds are thin and under such conditions the danger of icing usually is not serious.

The presence of a halo indicates clouds composed of ice spicules which, as previously stated, do not form any appreciable ice deposit.

Low pressure areas usually are more favorable for icing conditions than high-pressure areas, since the former are generally attended by considerable cloudiness and precipitation. Favorable icing conditions are likely to obtain in regions to the leeward of large bodies of water where temperatures of freezing, or lower, frequently occur; also over high terrain where flights at high elevations are necessary.

In closing, it is desired to state that one of the chief difficulties in a study of this kind is the frequent impossibility for the pilot or observer to classify correctly the type of ice formation since it usually is melted by the time the airplane reaches the ground. Since many of these flights were made before daylight this difficulty was especially pronounced. Also, there is a certain amount of confusion in the minds of many as to what constitutes rime and what clear ice. It is hoped that the descriptions given here will make possible a more accurate classification in this respect in future observations. It is believed, however, that so far as averages are concerned, the values found would not change appreciably with additional observational data.

It is desired to acknowledge the cooperation of the National Air Transport, Inc., Chicago, Ill., with the

Weather Bureau in the procurement of a number of airplane observations during ice-forming conditions. During the winters of 1928-29 and 1929-30, regular mail planes flying between Chicago and New York and between Chicago and Kansas City, were equipped with aero-meteorographs when conditions appeared favorable for ice formation. Local flights such as are now made daily at the Weather Bureau airport stations at Atlanta, Ga., Chicago, Ill., Cleveland, Ohio, Dallas, Texas, and Omaha, Nebr., however, provide far more satisfactory data for a study of this kind than do flights made over great horizontal distances.

I am indebted to Mr. L. P. Harrison of the Aerological Division, Weather Bureau, for many helpful suggestions during the preparation of this paper.

Weather Bureau,
Washington, D. C., October 26, 1932.

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^{*}A paper by W. Bleeker, entitled "Einige Bemerkungen über Eisansatz an Flugzeugen," Met. Zeit., Bd. 49, Heft 9, 1932, pp. 349-354, came to the attention of the author after the completion of this paper. Bleeker emphasizes the effect of evaporative cooling in the freezing of water droplets; also the time required for a drop to freeze ("Vereisungszeit") is considered to be proportional to the radius of the drop and thus a factor in determining the amount of deposit. Applications of the above, computations based thereon, and other topics are also discussed.

		7.1	BLE I.	FLI	ORT O	BERVA	IONS 1	(ADE DI	URING	THE POPOLATION OF ICE ON THE AIRPLANE, CLASS	SIFIED ACCORDING TO THE TYPE AND THIORIESS OF	PORMATION			
	Oloud or other conditions in which toe formed	Elevation above ground where ice began to form	Elevation above ground where ion desired to form	Temperature where	Temperature where to tors	Relative boundity where toe began to form	Relative bunddity Where for ceased	Time interval in which ice formed	Jo 222	. Parts of sirplane where ice formed.	Remarks	Place of ob	and d	late ;ion	
		Yators	Keter	90	00	Per-	Per-	mi n	18.	OLZAR IGE, 1/4 INGH	OR MORE				
91. A8t. A8t.	1	250 1,200 3,100	590 2,060 3,960	- 5	- 1 - 6 - 3	95 91 95	95 89 95	3 1/8 3 14	1/2 1/4 1/4	Wings, wires and struts All leading edges Wirs, struts and t.alling edges of wings	Smooth Yery plear and rather rough, seewing from	Oleveland, Ohiomgo,	Jen.	a,	1932 1932
ASt. StOu. StOu. StOu. StOu. StOu. StOu.	and rain	2,640 900 430 3,380 3,100 550 9,600	2,970 1,420 980 8,960 3,500 1,040 4,660	- 7 - 4 0	- 5 -10 - 2 - 7 - 8 - 3 -35	98 84 91 91 91 90	99	7 5 6 9 1 1/2 4	1/2 1/3 1/4 1/3 1/4 1/4 1/4	All leading edges Propeller, wings, wires and struts Leading edges all exposed surfaces Management of the surfaces Mings, wires and struts All leading edges Leading edges, all exposed parts	these clouds Forced to descend because of ice forsation Rough, 1/2 in. on wires, 1/4 in. on wings Flaky Sufficient to cause end of flight Rough Smooth, ice on wings made further climb impossible	Chicago, Chicago, Cleveland, Chicago, Chicago, Cleveland, Omaha,	Jan.	81, 84, 87, 5, 17,	1932 1932 1932 1932 1932
st0a.		4,140	4,450	L_	-15	90	98	4	1/4	Leading edges, all exposed parts	Smooth. Atrilans became so heavy with ice that it began to lose altitude and further climb impossible	. Omene.,	Bat.	19.	
8tOu. A6t. A8t. A8t. StOu.	and rain	850 1,420 3,870 4,450 1,870	1,370 1,900 4,370 4,640 1,630	- 5 - 2 -14	-10 - 8 - 8 -17 - 8	90 85 95 78	94	3 13 13 8	1/4 1/4 1/4 1/4	Leading edges, all exposed parts Wings, wires, struks and tail surfaces Leading edges Leading edges and propaller	Airplane wibrated and flight had to be	Chicago, Chicago, Chicago, Chicago,	Mar. Apr. Apr.	24, 24, 26,	1932 1932 1933
lt,-Çu.	talm bas	2,380	3,110	l l	- 5	95	95.	8	1/3	Wings, wires and strute	abandoned Rough, 1/3 in. on wires and strute, 1/4 in. on wings	Obloago, Oleveland,	Apr.		
A5t. A9t.	and rain	3,800 3,070	3,930 3,400	- 3 - 1	- 5 - 4	98 98	94 98	B 4	3/8 1/3	Wings, wires and strate Wings, wires and strate	Rough Rough, flight stopped because of ice	Oleveland, Oleveland,	31000	L.	LYM
										CLEAR IOE, LESS TH	A H 1/4 I H O H				
A5\$. 5\$0u.		2,859 1,100	8,890 3,860	- a	- 9	100 98	96	1 15	1/8 1/8	Wires and strute Leading edges, strute and Andahiold	Rough Smooth, airplane became so heavy with ice that pilot was mable to climb higher, much los resulted on airplane after landing	Oleveland,	May		1933
8tOu.		1,250	1,580	- 7	- в	83	67	1	Thin	On mero-metacrograph and various parts of mirplans	,	Ohiongo,	Mar.	19	1072
3t0u. 3t0u. StCu.		750 1,550 770	1,350 3,860 820 1,380	- 2	- 6 - 8 - 9	84 96 95	96 97 95	3 7 8	Thin 1/8 Thin	On more-meteorograph and leading edges leading edges and struts Leading edges of wings	Airplane seemed heavy at top of clouds	Omehe, Omehe, Omehe,	Hat. Jan. Jan.	an.	1933 1933 1933
it.—Cn.		1,170 570	1,240	- 5	- 8	86	84	2 1/2	1/8	Leading edges of wings	Hard organization ice, made very noticeable difference in the handling of simplene	Omelia.	Dec.	Б.	1933
tOu. ASt.	ınd AGu,	760 4,930 3,190	1,090 5,170 3,250 3,640	-17 -18	- 5 -18 -18	97 70 86	96 80 90	B 1 3	1/8 Thin Small	Leading edges of wings Windshield Windshield and leading edges	Frevented further amoent Hard crystal, smooth	Omaha, Omaha, Omaha,	Dec. Nov. Dec.	ai,	1931 1931 1931
81.		3,530 1,380	8,220	- 4	- 8	88	96	5	1/52	Leading edges of wings, struts and sero- meteorograph	Smooth, very noticeable as wires began to	Ossiba,	Feb.	,	1933
91. 8t.	,	950 840	1,180 1,180	- 1	- 6	96 90	96 95	1	Thin	Windshield	Otroaked in very thin patches Probably clear ice but insufficient to determine definitely	Chicago,	Dec.	жэ,	1931
pose at		1,700		-1	-	90	=		Small Mount	Wings, wires and atrute	Moisture carried from clouds from after smorging from clouds, fromen in drops	Oleveland,			1931
	alow St	1,100		- 3	- 2	84 85	-	1 1	Thin	Strute Trailing edge of strute, wings and on	Smooth, none formed in clouds Formed from moisture left on mirplane	Omaha,	Cat.	20,	1937
		900		+1	,	83	_		Thin	windshield. Trailing edge of wings	from clouds Formed from moisture left on givplens	Onicego,	0a \$.	31,	1931
POAS BI	uą.	4,150		۲,	l ']		heet	statement and and ands	from clouds	Dellas,	KAY	39.	1932

Olow or other conditions in which ice forms	Elevation above ground where toe began to form	Elevation above ground where ice ceased to form		Temperate	Relative homisty where toe began f to form	Relative bunidity where toe censed to form	Time interval in which ice formed	Thiolness of coat-	Parte of airplane where ion formed	Remarks	Place and : te of observation				
	Ketera	Katare		οσ	cent	oemt	min.	in.	RIMB, 1/4 INOH OR MORE						
510u. 510u. 510u. 510u. A0u.	1,390 1,300 730 2,250 3,590	1,690 1,600 990 8,010 4,000	- 4 - 5 - 5	- 3 - 8 - 5 - 7	96 85 92 90 85	96 93 98 94 96	5 4 8 88	5/18 1/4 1/4 3/8 3/4	Wings, wires and strute Wings, wires and strute Leading edges, atruts and acrossteorograph All surfaces 1/2 in. on ware, 3/4 in. on leading edges and wings	Hough Frosty white Airplane suddenly stopped olimb- ing due to ice, very rough and white	Oleveland, May 38 Oleveland, Jan. 27 Ohicago, Dec. 5 Ohicago, Feb. 16	, 1978 , 1931			
AGu. AGu. AB\$.	4,150 5,480 5,270	4,830 3,710 3,870	-10 -18 0	-14 -13 - 3	85 75 95	93 96 95	25 10 4	3/8 -	All over wirgs, struts and rigging All surfaces iroluding under side of wings Soticed on shield only	white Unable to clish higher due to loc Could not clish higher due to loc Probably formed on wings also as airplane becare very unmanage- able	Chicago, Dac. 82 Dallas, Dac. 30 Chicago, Feb. 9	, 1931 , 1932			
A81. A8t. A8t. A8t.	3,390 1,830 2,100 3,850 4,380	3,700 8,850 8,400 4,080 4,590	ı	-14	95 95 95	94 96 96 90	5 3 18	1/4 1/4 1/4 1/4	Wires and strute Wings, wires and strute Wings, wires and strute Wings, wires, strute and propeller	Frosty white	Oleveland, Nov. 12 Cleveland, Dec. 6 Cleveland, Dec. 9 Oleveland, Jan. 5	1931 1931 1932			
5%. 5%. 5%. 5%.	840 840 860 880 880	1,070 1,190 1,190 1,130 940	- 5 - 4 - 5	-10 - 6 - 9	96 95 96 96	93 98 97 100 100	5 1/2 5 1/2 4 1/2 5 2	1/4 3/6 1/2* 1/4 1/8	Wings, wires and strute Wings, wires and strute (Nore than? Wings, wires and strute Wings, wires and strute Wings, wires and strute	Ice formed rapidly and came off as	Oleveland, Jan. 35 Oleveland, Jan. 30 Oleveland, Dec. 35 Oleveland, Dec. 15	, 1931 , 1931			
St. St. St. St. St. StCu. rain and	960 460 500 870 660 3,450	1,120 1,020 940 1,470 980 4,840	- 7 - 3 - 5 - 3	- 8 - 3 - 4 - 8 - 5 -10	96 90 89 80 98 93	98 93 93 85 100 90	6 4 2 2 2 1/2 28	1/3 3/8 1/4 1/4 1/4	Wings, wires and strute Wings, wires and strute Wings, wires and strute Wings, wires and strute Wings, wires and strute Leading adgres wings and strute	Evaporated quickly in the clear air	Cleveland, Nov. 25 Cleveland, Feb. 3	, 1931 , 1931 , 1931 , 1932			
StOu. and rain StOu. and snow StCu. and snow St. and snow St. and rain St. and ASt.	5,090 640 680 980 1,260 980 5,060	3,630 1,130 1,800 1,760 1,890 1,140 5,180	0 -10 -10 - 5 - 0 - 13 - 7	-10 - 2	95 86 85 90 93 94	98 90 90 93 100 94	19 3 10 7 3 1	1/2 1/4 1/4 8/8 1/3 1/4 -	Leading edges wings and struts Whige, wires and strute Whige, wires and strute Leading edges wings, wires and strute Whige, wires and strute Whige, wires and strute		Dallas, Jan. 4 Uleveland, Feb. 18 Gleveland, Feb. 25 Gleveland, Dec. 1 Gleveland, Dec. 4 Gleveland, Dec. 10	1952 1952 1951 1951 1931			
ASt. and snow ASt. rain and sleet ASt. rin and sleet	4,000 4,080 3,480	4,820 4,410 4,800	- 4	-11	90 90 94	92	3 1/2 8 1/3 -	1/4	Wings, wires and struts Wings, wires and struts Leading edges, struts and all rigging	"Yood so heavily had difficulty allabing, white heeds of ics	Gleveland, Jan. 21 Gleveland, Nov. 14 Dallas, Dec. 8	, 1931 , 1931			
ASt. and snow ASt.	2,680 2,670 560	3,390 3,350	- 6 - 2	-13	90 94	93 92 95	34 4 5	1/4 5/8 1/4 3/8	Wings, wires and strute Wings, wires and strute Wings, wires and strute	loe uneven on wires and struts,	Oleveland, Feb. 31 Oleveland, Mar. 2 Oleveland, Mar. 6	, 1933 , 1933			
St. StOu. StOu. StOu. ASt. and snow Stou. Stand snow St. and snow ASt. and rain	810 540 1,200 430 1,180 2,480 560 700 930 2,670	1,500 1,430 1,540 1,590 2,680 1,180 1,430 2,300 4,950	- 4 - 8 - 7 + 1	-13 -6 -10 -8 -0 -5 -10	96 95 90 94 98 97 90 98 95	96 99 95 95 95 96 98 100 95 90 88	2 1/8 5 4 3	3/9 1/4 1/4 1/8 1/8 1/4 1/2 1/2	Wings, wires and strute Unings, wires and strute Unings, wires and strute Unings, wires and strute Of wings	White frosty Rough *About one dozen icicles, yellowish rise	Oleveland, Mar. 10 Oleveland, Mar. 19 Oleveland, Mar. 19 Oleveland, Mar. 32 Oleveland, Jan. 27 Oleveland, Jan. 27 Oleveland, Apr. 13 Oleveland, Apr. 13 Oleveland, Apr. 13	, 1932 , 1932 , 1932 , 1933 , 1952 , 1933 , 1938 , 1938			
A0u. St0u. A0u.	4 200 850 2,350	4,450 1,450 8,550	-11 - 2 - 3	- 4	70 94 93	90 95 93	6 5	1/4 5/8	Wings, wires and struce Wings, wires and struce	, June Water Land	Cleveland, Apr. 36 Cleveland, Apr. 38	, 1952			

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506	ಕ್ಷಿಕ್ಷ್ಮ	5 5 =	to form	48	1 d d	humidity occsed	44	100	OLASSIVIED ADDOB	NOTATION SSERVICE OF JAKE BALL DATA	
Oloud or other sonditions in which ice formed	Elevation above ground where toe began to form	Elevation above ground where tos commed to form	Temperature 10e began to	Teaperature where	Relative humidity where ice begins to form	Relative hum where toe on to form	Time interval	Thiokness of ing	Parts of sixplame where ice formed	Homeska	Place and date of observation
	Mutora	Meters	93	ۍ	Per-	Per-	mln.	in.	RINE, LESS THAN	1/4 INCH	
AOu. AOt. StOu. St. St.	3,510 4,420 2,370 450 430 680	3,870 4,570 2,750 750 940 2,150	- 3 -13 - 3 - 3 - 1	-4 -14 -5 -4 -3	95 78 94 90 94 95	85 88 85 95 97	7	1/8 1/15 1/16 1/8 Trace Small	Wires Leading edges Leading edges Hings, wires and struts Windshield Leading edges	Frosty, white Whitisch Whitisch and escoth Light most at ground but some noted by tilet in clouds	Oleveland, June 5, 1932 Chiosgo, May 9, 1932 Chiosgo, May 10, 1932 Cleveland, Jun. 4, 1932 Cumha, Dec. 34, 1931 Cumha, Feb. 14, 1952
8t. 8t. AOu.	ļ	1,150 960 5,150	- 9 -18	-10 -1 -20	95 96 88	98 97 82	1 15	1/8 1/16 Very thin	Wings, wires and strate Leading adget Wings and strate	Heather airplane or instrument showed any vibration	Oleveland, Peb. 5, 1933 Ohicago, Peb. 2, 1932 Ballna, Jan. 18, 1932
StCu, and ecow St. and rain	4,040	980 4,500	-10 - 8	-12 - 3	94	94	4	7hin 1/16	Strute Leading edges wings and struke	Formed immediately upon extering clouds and sulted immediately upon learning clouds los disappeared on descent at 3,500 meters	Chiongo, Nov. 25, 1951
St. and rain AOu. StOu. ASt. StOu.	3,280 940 3,490 1,250	1,550 3,610 1,930 3,960 1,600	+ 1 -18 - 6 -13 - 9	- 6 -20 - 8 -17 -12	94 95 95	95 100 88 97 93		1/8 Thin 1/8 1/8 Small	Wires and strate Leading edges exposed parts Hings, whree and strate Wings, whree and strate Wires	shows ground Smooth, not noticed until after landing Frosty, white	Ohtoago, Oct. 15, 1951 Claveland, Cot. 30, 1951 Csaha, Har. 8, 1953 Cloveland, Har. 1, 1953 Cloveland, Har. 16, 1953 Claveland, Har. 34, 1953
56. 810u.	400 1,740	590 4,750	- 1	- 2 -19	95 95	98 1.00	8	1/3	Wings, wires and strute Lending edges exposed parts	Not noticed during flight, seen after landing	Cleveland, Mar. 89, 1952 Camba, Jan. 14, 1952
A02. A65. A04.	4,090 3,400 2,400	4,140 4,170 2,780	-10 -14	-31 - 9 -17	58 93 60	70 100 90	8 7 4	1/8 1/8 1/8	Leading edges and wader side of wings Leading edges and wings Leading edges and landing goar, conding and aerometeorograph		Oniongo, Mar. 15, 1932 Oniongo, Apr. 6, 1932
60. and rais AOu.	1,100	1,550 5,430	+11	- 1 -18	99 70	100	3 18	1/8 Small motest	Wings, whree and strute	Rough Frosty	Ohloseo, Mar. 38, 1932 Oleveland, Apr. 11, 1932 Omaha, Apr. 28, 1933
ACu. StCu.	390 430	5,390 1,040 1,290 3,650 610 810 1,020 980 1,300 470 680	- 9 + 1 - 7 - 7 - 7 - 7 - 7 - 7 - 7 - 7 - 7 - 7	- 8 - 8 -10 - 5	62 90 90 90 92 83 90 90 95 95	90 95 65	1 1/8 2 1/8 3 1/8 1 1/8 1 1/8	1/8 3/16 3/16 1/8 1/8	Wings, wires, struts Drace wires only Brace wires only Wings, whree and struts Wings, whree and struts Wires Brace wires only	White, frosty Frosty white Prosty white Flaky	Oleveland, Out. 23, 1931 Oleveland, Out. 17, 1931 Oleveland, Jan. 19, 1933 Oleveland, Jen. 2, 1933 Officego, Jan. 8, 1932 Officego, Jan. 29, 1932 Officego, Jan. 29, 1932 Oleveland, Feb. 4, 1932 Oleveland, Feb. 28, 1932 Officego, Jan. 24, 1932 Ohicago, Jan. 24, 1932
810u. 810u. 810u. 810u.	1,950 720 550 640	1,530 1,070 910 930	- 4 - 2 - 5 - 3	- 7 - 5 - 7 - 6	88 84 89 84	94 94 95	-	1/16 Light 1/16 Thin	Leading edges, strute and wires Leading edges and wings Leading edges Leading edges		Chicago, Nov. 8, 1931 Jhicago, Rec. 25, 1931 Chicago, Feb. 5, 1932 Chicago, Feb. 24, 1932
ASt, Slee and anow AOu. AOu. AOu.	2,600 4,860 4,110 3,550	4,150 5,000 4,190 5,810	- 5 - 8 - 4 -11	-18 - 9 - 4 -15	92 85 96 100	87 95 96 100	35	1/8 1/15 1/8 Thin	' • ' - • •	los disappeared on descent at 1,500 meters	Oleveland, Apr. 10, 1938 Ohicago, Oct. 8, 1931 Ohicago, Oct. 85, 1951 Dallas, Feb. 24, 1932
ACt. ACt. ACt. ASt.	4.350	4,900 8,990 4,490 3,140	-18 -15 -19 -10	-30 -18 -91 -12	85 72 90 90	86 74 90 96		1/8 Thin 1/16 Small	Wings, wires and strute Leading edges Wires Wires and strute	Ground lights visible through clouds	Cleveland, Feb. 35, 1938 Chicago, Feb. 1, 1933 Cleveland, Feb. 37, 1933 Cleveland, Feb. 15, 1938
A9t. A0u. BtOu.	4,490 3,640 4,850	4,910 3,790 5,060	-13 -11	-17 - 8 -15	84 97 63	88 94 70	7 2 2	1/8 1/16 6==11	Leading edges Viros Leading edges	Soft, made airplane bard to headle	Omaha, Doc. 8, 1981 Cleveland, May, 7, 1982 Omaha, May 25, 1982
ASt. StCu.	3,490 3,770	4,440	- 9	- 8 -14	88 72	95 40		1/39 1/8	Strute, wires and leading edges, wings Leading edges strute and wires	Rough	Omeha, May 24, 1958 Chicago, Apr. 21, 1958

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Cloud or other conditions in which toe formed	Elsvation above ground where tos began to form	Elevation above ground where tos ceased to form	Temperature where	Tampersture where	Relative hunidity where toe began to form	Relative humidity where toe cessed to form	anter	Intomess of cont-	Parts of simpleme where ice formed	Romarks	Place of ol	and date
	Metore	Motors	°О	°a	Per-	Per-	win.	in.	CLEAR ION, TELCENSS	U # X # O # #		
Mist and	1,900	3,500	- 1	-10	70	90	1,8	2	leading edges and asroneteorograph	Small globules of clear toe	Omaha,	Feb. 25, 1932
A8t.	2,350	2,760	a	- 2	91.	89	8	?	All over girplane	Clear drops, much moisture on nirplane before ice formed	Chicago,	Dec. 30, 1931
Rain and StOu.	800	1.060	0	_ 1	88	90	5	7	Wires and metal-parts, not on wings	Smooth	Omelta,	Oct. 19, 1931
##Ou, Rain	550	730	- 3	- 4	84	86	8	3	Lending adges	Rough	Onlongo,	Jan. 15, 1932
(Below	4,220	5,000	- 3	- 7	98	83	20	7	Wires and leading edges, wings and struts	About the size of raindrops	Dallas,	Jan. 18, 1932
				<u> </u>	_		-		RIME, THIORNESS UNK	HOWN	 	
StOu. rain and sleet	5,150	3,950	0	- 8	100	80	. 6	3	Leading edges wings, struts and wires	Like small raindrops, white	Dallas,	Feb. 15, 1952
BtCu.	3,610	4,440	_ 3	- 7	3.00	97	13	,	Tings, struts, rigging and serossteorograph	White bonds -	Dallas,	Feb. 16, 1952
StOn.	8,400	2,800	0	- 3	100	98	8	?	Leading edges, strute and wings	Like frozen raindrope, white	Dallas,	Feb. 22, 1952
8tOu. rain	3,590	4,610	_ 1	- 8	99	94	25	7	Strute, wires and trailing edge of wings	Whitish, about size of rain-		
and sleet ASt. ASt.	4,030	4,330	- 9	-18	85	95	8	9	Struts and bottom of wings	drops, flattened into flakes Smooth	Dalles, Chicago,	Jan. 28, 1933 Dec. 31, 1931
rain and elect	3,120	4,840	- 8	-12	82	93	24	7	Leading edges wings, strute and rigging wires	Smooth	Dallas,	Dec. 1, 1931
									RINE, AND OLEAR ICE IN BA	ME OBSERVATION		
AOu.	4,400	4,700	- 3	- 5	62	85	5 1/3	174	Wings, wires and strate	Inner coat frosty, outer coat	(I) errel err	l Dot. 24, 1931
9t0u.	630	1,160	- 6	-10	B≋	95	8		beading edges	Smooth, older and molid, except		
				ļ				1		rough whitish strip center o	f	
						Į.				formerd edge of wing, adherin	rg I	
				}	}		}			very tightly to airplane,	}	
							ļ			thickest at exact center of and leading edge and distints	 .x_	
		1		ĺ			ľ			ing within several inches of	1	
8tCu.	1,820	2,600	- 7	-18	84	95	 - 	1/3	Wings, wires and struts	leading edge Semiclear ice and memifrosty	Ohioago, Olevelan	Dec. 6, 1951 1, Nov. 2, 1951
BtCa.	3,290	4,540	-1	- 9	98	98	i '	i Hory	Loading odgos, struts, wings and wires	Snow frome first as whitish	[•

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TABLE II.

				_	_				
Type	Thick-	No. of	Average	Average	Average	Average	Average	Average	Average
	ness	cases	temper-	temper-	relative	relative	time	height	thickness
]		ature at	ature	humidity	humidity	during	above	of stratum
į			which	interval	at which	interval	which	ground	în which
	}]	formation	in which	formation	in which	formation	where	formation
	-		occurred	formation	occurred	formation	occurred	formation	occurred
•				occurred		occurred.	[occurred	
•	j					(+ indicates]		
1	1					higher hu-			
						midity at	1		
	ļ					end of for-	}		
	ļ					mation than			
	[1		at beginning;	[
	j] [j	}	vice versa.)	!		
			c	°C	Percent	Percent	min.	meters	meters
Clear	1/4 in.								
ice	or more	19	-5.8	3.5	92	+ 2	7.5	2,476	576
Clear	Less	-			ļ		1	·	
ice	than			i -	1		i !		
	1/4 in.	19	-5. 9	2,4	90	+ 3	3.7	1,943	499
Clear	Un-	-		1	Ì			_	
ice	known	5	-3.1	3.8	86	+ 2	8.4	2,272	696
Clear			1			i	<u> </u>	-	
ice	All	43	-5.5	3.1	90	+ 2	6.3	2,245	562
				ļ	·		f [
Rime	1/4 in.				į.] !		
	or more		-6.3	3.3	92	+3	6.6	2,151	522
Rime	Less			1	İ		!	·	
_	than]		1	Ì				
	1/4 in.	52	-8.6	2.7	89	+2	4.8	2,482	395
Rime	Un-	Ì		•			!	-	
	known	6	-5.2	5.3	93	- 2	13.7	3,764	828
Rime	All	108	-7.3	3.1	91	+ 2	6.1	2,399	479
								,	
Clear	All	4	-6.6	4.8	87	+10	9.5	2,891	718
ice		-		,	1	İ		· •	1

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and rime during same observation

N.A.C.A. Technical Note No. 439

TABLE III.

Types	Thickness	N	i		
Clear ice	1/4 in or more Less than 1/4 in. Unknown	Chicago 10 · 3 2 15	Cleveland 6 2 0 8	Dallas O 1 1	Omaha 3 11 2 16
Rime Rime Rime Rime	1/4 in. or more Less than 1/4 in. Unknown All	4 20 1 25	38 22 0 60	4 2 5 11	1 8 0 9
Clear ice and rime during same observation		1	2	1	

TABLE IV.

	Average te ature dur ice forma as obtain data in T	ing the tions ed from	1931 to	or Nov. April, ncl.,	Average amount of low clouds. 8:00 a.m. E.S.T. for Nov., 1931 to April, 1932, incl. (Scale 0-10)	Average amount of intermediate clouds. 8:00 a.m., E.S.T., for Nov. 1931 to April, 1932, incl. (Scale 0-10)
****	Clear ice	Rime	2,245 meters*	2,399 meters*	*	
Chicago Cleveland Dallas Omaha	-4.8 -2.8 -3.0 -7.4	-8.1 -6.8 -7.1 -9.5	-2.6 -2.8 6.5 -0.5	-3.3 -3.4 4.9 -1.2	4.3 5.0 4.4 4.8	1.2 2.0 C.6 O.8

^{*} Average height above ground at which clear ice formed.

** Average height above ground at which rime formed. (See Table II.)

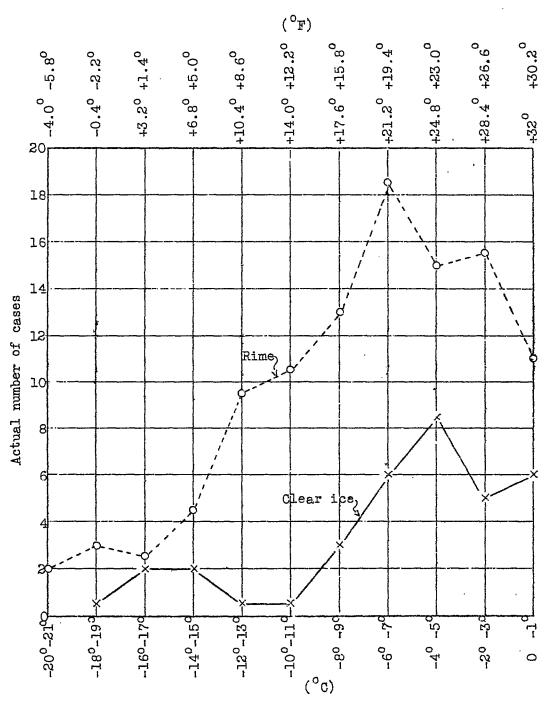


Fig.1 Actual number of cases when clear ice and rime formed at various temperatures, the latter representing the mean of the temperature range through which each formation occurred.

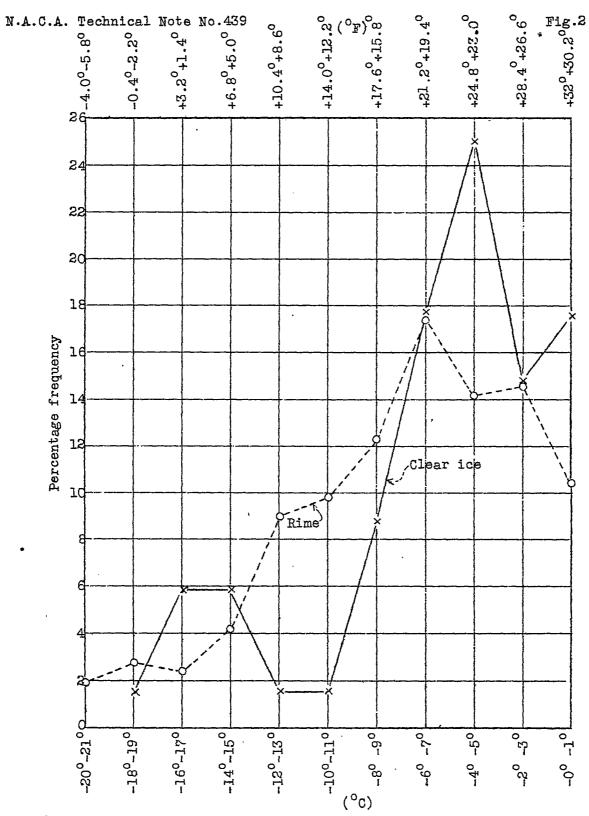


Fig.2 Percentage frequencies of clear ice and rime formations at various temperatures, the latter representing the mean of the temperature range through which each formation occurred

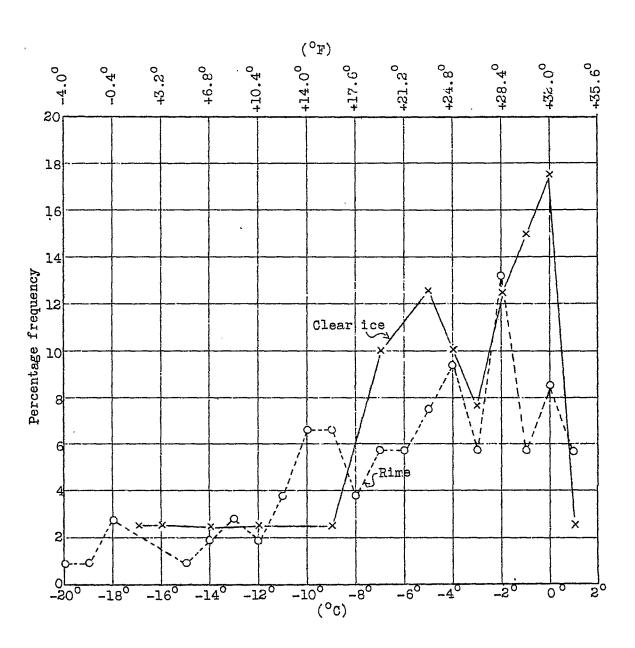


Fig.3 Percentage frequencies of clear ice and rime formation at the temperatures at which the ice began to form.

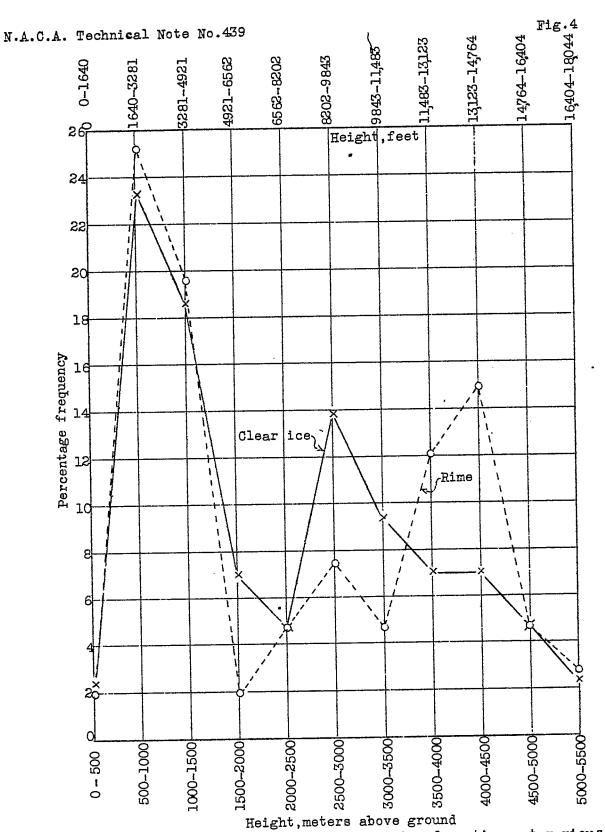


Fig. 4 Percentage frequencies of clear ice and rime formations at various heights above the ground, the latter representing the mean of the height interval in which the formations occurred.

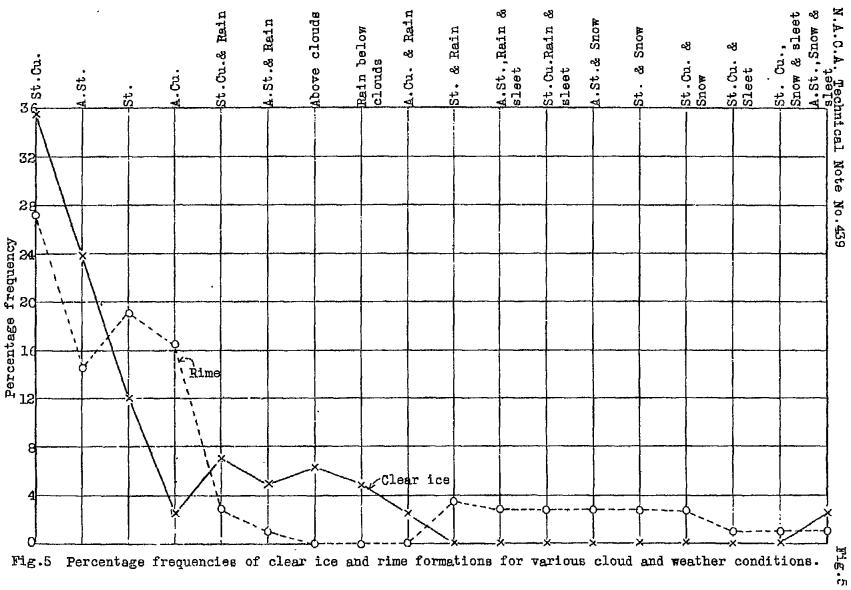


Fig.5 Percentage frequencies of clear ice and rime formations for various cloud and weather conditions.